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## SEMIANNUAL TECHNICAL REPORT **NORSAR PHASE 3**

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## VII.7 On the M<sub>s</sub>-m<sub>b</sub> Relationship of Earthquakes

The purpose of this study is to estimate the  $M_s-m_b$  relationship using data from VLPE, ALPA, NORSAR and PDE, and assess the impact of network magnitude bias effects (Ringdal, 1976) on the results. We assume that there exists a linear functional relationship between  $M_s$  and  $m_b$  of shallow earthquakes of the form

$$M_{c} = \alpha \cdot m_{b} + \beta + \varepsilon$$

where  $\alpha$  and  $\beta$  are unknown coefficients and  $\varepsilon$  is a random variable that is normally distributed with zero mean and an unknown variance. Generally, the formula (1) can only be used as an approximation over limited magnitude ranges; for example, Gutenberg and Richter (1956) estimate  $\alpha$ =1.6 using only very large events, while Evernden (1975) finds a value of  $\alpha$ =1.0 below magnitude m<sub>b</sub>=5.0 down to at least m<sub>b</sub>=3.0.

This study is restricted to the magnitude range of most interest for current seismic discrimination studies, i.e., approximately  $m_b=4.0$  to 6.0. The data base consists of 52 randomly selected shallow Eurasian earthquakes (Turnbull et al, 1975). Magnitudes of these events have been available from PDE and NORSAR ( $m_b$ ), and from VLPE, ALPA and NORSAR ( $M_s$ ). In addition, we used the maximum-likelihood technique of Ringdal (1976) to modify the PDE  $m_b$  estimates; we denote these modified estimates by PDE(m-1)  $m_b$  values. The assumptions required to obtain the PDE(m-1) estimates are described in detail by Turnbull et al (1975).

In order to examine the variations of the  $M_s - m_b$  slope as a function of estimation techniques of  $M_s$  and  $m_b$ , a total of nine cases were run based on the given event population. In each case, the M values estimated by either VLPE (averaging), ALPA or NORSAR were combined with the  $m_b$  values

(1)

of NORSAR, PDE and PDE(m-1). One case is shown in Fig. VII.7.1, while complete results are summarized in Table VII.7.1. The following points are noteworthy:

- 1. Four of the runs produce a virtually identical slope  $(\alpha \approx 1.37)$ . These are precisely those four runs that combine "consistent" M<sub>s</sub> and m<sub>b</sub> values, i.e., values free of network bias effects. (NORSAR or ALPA M<sub>s</sub> versus NORSAR or PDE(m-1) m<sub>b</sub>).
- 2. A consistently high value of the slope (1.66 or 1.64) is found when PDE  $m_b$  is plotted against a consistent  $M_c$ .
- 3. A consistently low value of  $\alpha$  (1.23 or 1.24) results when VLPE M<sub>s</sub> is plotted against a consistent m<sub>b</sub>.
- 4. When PDE  $m_b$  is plotted against VLPE  $M_s$ ,  $\alpha$  is again high, showing that the network bias effects in PDE magnitudes dominate those of VLPE.

Hence, the behavior of the computed slope agrees well with what could be expected from network bias considerations (network magnitude are expected to produce a bias that is largest for small events). It appears that the most accurate linear functional relationship between  $M_s$  and  $m_b$  for the given data set (ranging in  $m_b$  values from about 4.0 to 6.0) has a slope of approximately 1.4.

Considering more closely the four cases of consistent estimates, it is interesting to note that the value of the orthogonal standard deviation  $\sigma$  is lower when using PDE(m-1) m<sub>b</sub> versus either ALPA or NORSAR M<sub>s</sub> compared to when NORSAR m<sub>b</sub> is used ( $\sigma \approx 0.26$  vs  $\sigma \approx 0.31$ ). It would be interesting to compare PDE(m-1) m<sub>b</sub> to the VLPE-ALPA-NORSAR combined network with the network bias reduced by maximum likelihood processing; however, we have not been able to do this, mostly because of the lack of reliable VLPE data for some stations.

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## REFERENCES

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Gutenberg, B., and C.F. Richter (1956): Magnitude and energy of earthquakes, Ann. di Geofis., 9, 1-15.

Ringdal, F. (1976): Maximum-likelihood estimation of seismic magnitude, Bull. Seism. Soc. Am. (in press).

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## Table VII.7.1

VLPE

ALPA

VLPE

ALPA

VLPE

ALPA

NORSAR

NORSAR

NORSAR

PDE

PDE

Maximum

NORSAR

Likelihood

measurement procedures.					
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<sup>m</sup> b	M S	No. of Points* NP	Slope a	Intercept β	σ Orthogonal

1.54

1.66

1.64

1.23

1.37

1.37

1.24

1.37

1.39

-3.58

-4.33

-3.95

-1.73

-2.57

-2.25

-1.67

-2.48

-2.14

0.238

0.253

0.262

0.257

0.259

0.309

0.321

0.305

52

41

35

52

41

35

52

41

35

Estimated relationship M -m for various combinations of M and m b measurement procedures.

The missing data points (NP<52) are due to lack of available data for NORSAR or ALPA M for some events, and not due to nondetection at these stations.

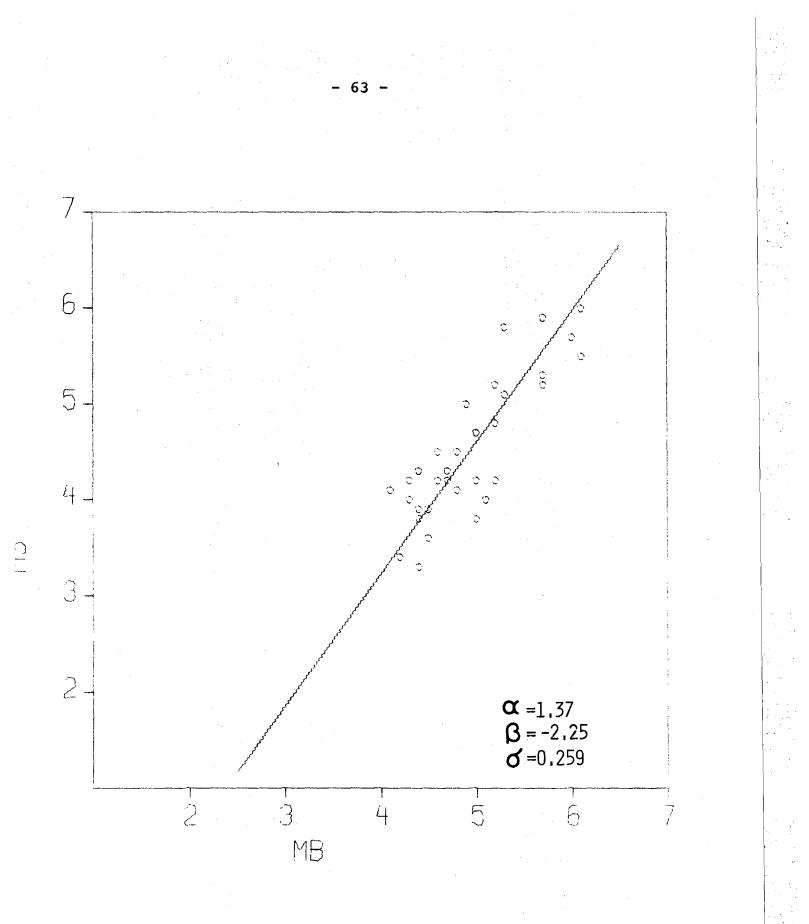


Fig. VII.7.1 PDE m modified by maximum likelihood processing plotted against NORSAR M s.