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## 6.3 Surface Wave Tomography for the Barents Sea and Surrounding Regions

### 6.3.1 Introduction

Existing global and regional tomographic models have limited resolution in the European Arctic due to the small number of seismic stations, relatively low regional seismicity, and limited knowledge of the crustal structure. During the last decades, new seismic stations were permanently or temporarily installed in and around this region. However, many of the data from these stations are not easily accessible via the international data centers but only by direct request to the different data operators.

Recently, a new crustal model of the Barents Sea and the surrounding areas had been derived in a joint project of the University of Oslo, NORSAR and the USGS (Bungum *et al.*, 2004; 2005). This model, with its detailed information on crustal thickness and sedimentary basins in the area, is a valuable constraint for the tomographic inversion of the upper mantle velocity structure based on surface wave data.

### 6.3.2 Data collection

In this project, we have extensively searched for long period and broadband data observed at the seismic stations and arrays in the area from the beginning of the 1970s until 2005. We were able to retrieve surface wave observations from the data archives at NORSAR, University of Bergen, the Kola Science Center in Apatity, the Geological Service of Denmark and the University of Helsinki, in addition to the data, retrieved from the international data centers IRIS and GEOFON. The full list of used stations is given in Table 6.3.1 and a map with the positions of the stations is shown on top of Fig. 6.3.1.

In these data archives, not yet analyzed Love- and Rayleigh-wave data were identified and for more than 150 seismic events (earthquakes and nuclear explosions) dispersion curves were measured. The list of these events is presented in the Table 6.3.2 and the map on bottom of Fig. 6.3.1 shows the geographic distribution of the events.

### 6.3.3 Data analysis

From the surface wave recordings, group velocities of Love and Rayleigh waves were measured in the period range between 10 and 150 s using the program package for Frequency-Time Analysis developed at the University of Colorado (Ritzwoller & Levshin, 1998). After several cleaning procedures (Ritzwoller & Levshin, 1998), the new measurements were combined with the existing set of group velocity measurements provided by the Center for Imaging the Earth's Interior at the Colorado University (CU; see Levshin *et al.*, 2001). From this CU data set only those paths were selected, which were completely inside the cell [50 – 90° N, 60° W – 160° E], such that the entire data set consists of paths within the same regional frame. From cluster analysis (Ritzwoller & Levshin, 1998) of the root-mean-square of the group velocity measurements for the new data set in the considered period range was estimated as 0.010 – 0.015 km/s for Rayleigh waves and 0.015 – 0.025 km/s for Love waves.

To demonstrate the amount of new surface wave observations, we compare in Fig. 6.3.2 the number of newly analyzed Love and Rayleigh waves with the number in the preselected CU

data set. Obviously, the new data set increased the ray density and consequently the resolution of the planned tomography. In particular for shorter periods, the number of rays crossing the target area was increased by more than 200% for Rayleigh waves and close to 200% for Love waves. For longer periods (*i.e.*,  $T > 80$  s), the ratio of added data significantly drops since large seismic events, necessary to generate long period radiation, are very rare in this region.

### 6.3.4 2D-inversion for group velocity tomographic maps

A tomographic inversion for two-dimensional group-velocity maps has been done for a set of periods between 14 and 90 s following the procedure by Barmin *et al.*, 2001. In all cases, we inverted the combined data set of the newly acquired and analyzed data and the preselected CU data. As first results, we present in Figs. 6.3.3, 6.3.4, and 6.3.5 the resulting group-velocity maps for Love and Rayleigh waves at three different periods: 16, 25, and 40 s, respectively.

To illustrate the newly achieved, high path density, we also present in Figs. 6.3.3, 6.3.4, and 6.3.5 all ray paths of the newly acquired data for both, Love and Rayleigh waves on top of the figures. In the middle of the figures, we show these ray paths again but with all rays paths of the preselected CU data set on top. Note the many gaps in the preselected CU data set, for which group velocity information is now available.

The Love and Rayleigh group-velocity maps are shown at the bottom of Figs. 6.3.3, 6.3.4, and 6.3.5. The group-velocity maps derived from the combined data set, show the lateral deviation of the group velocities from the average velocity in percent. Note that these deviations are up to  $\pm 36\%$  for Love and Rayleigh waves with a period of 16 s. This reflects the strong lateral heterogeneity of the Earth's crust in this region, which changes between the mid-oceanic ridge system (white line on the map), thick sedimentary basins in the Barents Sea, and old continental shields.

Sensitivity kernels for Love waves and Rayleigh waves are different; Rayleigh waves are sensitive for deeper structures than Love waves with the same signal period. Therefore, jointly analyzing the crustal structure with both Love and Rayleigh waves gives additional confirmation for an inverted velocity model. Note, *e.g.*, that the geographical pattern of group-velocity variations for Love waves at a period of 25 s (Fig. 6.3.4, on the right at the bottom) is similar to the pattern of group-velocity variations for Rayleigh waves at a period of 16 s (Fig. 6.3.3, on the left at the bottom). A similar relation is between the 40 s Love waves and the 25 s Rayleigh waves.

### 6.3.5 Further work: 3D-inversion for the shear velocity structure

The planned next step will be an inversion of all group-velocity maps (Love and Rayleigh waves) into a 3D shear-velocity model for the whole region (Shapiro & Ritzwoller, 2002). As additional constraints for the inversion the thickness of the sedimentary layers and the Moho depth, as they were derived by Bungum *et al.* (2004; 2005) will be included. The result should yield a new, robust 3D model of the velocities in the upper mantle beneath the greater Barents Sea region down to about 250 km.

### Acknowledgements

For this study we requested and retrieved broadband and long periodic data from the Norwegian National Seismic Network (NNSN, University in Bergen), Kola Regional Seismological Center (KRSC, Apatity), Danmarks og Grønlands Geologiske Undersøgelse (GEUS, Copenhagen), Totalförsvarets forskningsinstitut (FOI, Stockholm), British Geological Service (BGS, Edinburgh), the Finnish National Seismic Network (FNSN, University of Helsinki), and the international network operators and data centers IRIS and GEOFON. That they all made their data available for this study is gratefully acknowledged.

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**Christian Weidle, University of Oslo**

**Johannes Schweitzer**

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**Table 6.3.1. List of seismic stations from which surface-wave data could be retrieved to increase the ray coverage in the Barents Sea and surrounding regions (see also Fig. 6.3.1). The abbreviations in the network-affiliation column stand for: AWI – Alfred-Wegener-Institute for Polar and Marine Research, BGS – British Geological Service, CNSN – Canadian National Seismograph Network, FNSN – Finish National Network (University in Helsinki), FOI – Totalförsvarets forskningsinstitut (Sweden), GEUS - Danmarks og Grønlands Geologiske Undersøgelse, GRSN – German Regional Seismic Network, IDA – International Deployment of Accelerometers, IMS – International Monitoring System, IRIS – Incorporated Research Institutions for Seismology, KRSC – Kola Regional Seismological Center, MASI99 – Temporal net of seismic stations in Finnmark operated by NORSAR and the University of Potsdam (Germany), NNSN – Norwegian National Seismic Network (University in Bergen), RUB – Ruhr University Bochum, UK – University of Kiel, and USGS – US Geological Survey. The data availability for some of the stations may be longer than known to us.**

Station	Latitude	Longitude	Location	Network Affiliation	LP or BB Data Availability
AMD	69.7420	61.6550	Amderma	KRSC	08.1998 -12.2003
APZ9	67.5686	33.4050	Apatity	KRSC	09.1992 -
ARE0	69.5349	25.5058	ARCES Array	NORSAR/IMS	09.1987 -
ARU	56.4302	58.5625	Arti	IRIS/IDA/IMS	09.1989 -
BER	60.3870	5.3348	Bergen	NNSN	01.2004 -
BILL	68.0651	166.4524	Bilibino	IRIS/USGS/IMS	08.1995 -
BJO1	74.5023	18.9988	Bjørn Øya (Bear Island)	NNSN	06.1996 -
BORG	64.7474	-21.3268	Borgarnes	IRIS/IDA	07.1994 -
BSD	55.1139	14.9147	Bornholm	GEUS	01.1996 -
COP	55.6853	12.4325	København	GEUS	09.1999 -
DAG	76.7713	-18.6550	Danmarkshavn	GEUS/GEOFON/AMI	06.1998 -
DSB	53.2452	-6.3762	Dublin	GEOFON	12.1993 -
EDI	55.9233	-3.1861	Edinburgh	BGS	06.1996 -
ESK	55.3167	-3.2050	Eskdalemuir	IRIS/IDA/BGS	09.1978 -
FIA1	61.4444	26.0793	FINES Array	FNSN/IMS	04.2001 -
HFC2	60.1335	13.6945	Hagfors Array	FOI/IMS	08.2001 -
HFSC2	60.1326	13.6958	Hagfors Array (closed)	FOI	01.1992 - 09.2003
HLG	54.1847	7.8839	Helgoland	UK/GEOFON	12.2001 -
IBBN	52.3072	7.7566	Ibbenbüren	RUB/GEOFON/GRSN	07.1999 -
JMI	70.9283	-8.7308	Jan Mayen (closed)	NNSN	10.1994 - 04.2004
JMIC	70.9866	-8.5057	Jan Mayen	NORSAR/IMS	10.2003 -
KBS	78.9256	11.9417	Ny-Ålesund	NNSN/AWI/GEOFON/ IRIS/USGS	10.1986 - 09.1987 11.1994 -
KEV	69.7553	27.0067	Kevo	FNSN/IRIS/USGS	10.1981 -
KIEV	50.6944	29.2083	Kiev	IRIS/USGS	01.1995 -
KONO	59.6491	9.5982	Kongsberg	NNSN/IRIS/USGS	09.1978 -
KWP	49.6305	22.7078	Kalwaria Paclawska	GEOFON	06.1999 -
LID	54.5481	13.3664	Liddow	GRSN/GEOFON	01.1994 - 11.1995
LRW	60.1360	-1.1779	Lerwick	BGS	08.2003 -
LVZ	67.8979	34.6514	Lovozero	IRIS/IDA	10.1992 -
MA00	69.5346	25.5056	at ARCES A0	MASI99	08.1999 - 10.1999
MA01	69.3752	24.2122	Suosjavre	MASI99	05.1999 - 10.1999
MA02	69.1875	25.7033	Kleppe	MASI99	05.1999 - 10.1999
MA03	70.0210	27.3962	Sirma	MASI99	05.1999 - 10.1999
MA04	69.7127	29.5058	Neiden	MASI99	05.1999 - 10.1999
MA05	69.4533	30.0391	Svanvik	MASI99	05.1999 - 08.1999
MA06	70.4813	25.0609	Russenes	MASI99	05.1999 - 10.1999
MA07	69.7050	23.8203	Sautso	MASI99	05.1999 - 10.1999
MA08	70.1278	23.3736	Leirbotn	MASI99	05.1999 - 10.1999
MA09	69.4566	21.5333	Reisadalen	MASI99	05.1999 - 10.1999
MA10	69.5875	23.5273	Suolovuobme	MASI99	05.1999 - 10.1999



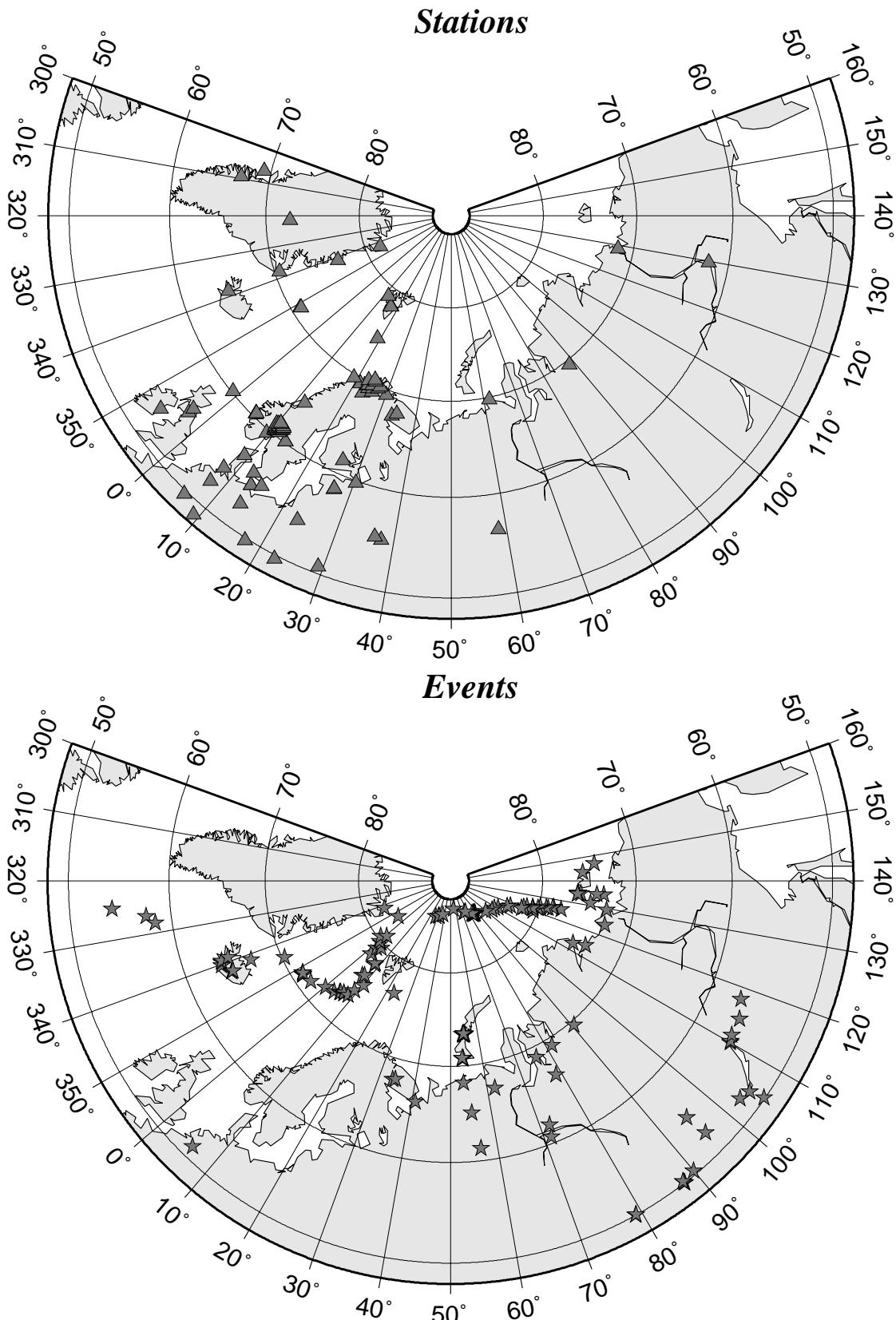


Fig. 6.3.1. The top map shows the seismic stations listed at Table 6.3.1 and the bottom map shows the distribution of the newly investigated seismic events as listed at Table 6.3.2.





Lat	Lon	Depth	Year	Month	Day	Time	mb	Ms	comment
79.5140	4.1880	10.0	2001	7	16	14:09:29.240	5.0	4.5	
80.8480	0.7680	10.0	2001	12	8	06:44:22.020	5.1	4.8	
85.8580	27.6810	10.0	2002	5	3	11:19:20.320	4.1	-	
86.0050	31.5950	10.0	2002	5	3	11:20:51.540	5.2	5.4	
85.9720	31.1490	10.0	2002	5	3	15:33:34.880	5.1	5.1	
86.2760	37.1770	10.0	2002	5	28	15:39:01.550	4.9	4.7	
75.6340	143.7460	10.0	2002	6	4	00:05:07.170	4.8	-	
84.0830	110.6900	10.0	2002	6	9	21:20:38.750	4.5	-	
83.1360	-6.0780	10.0	2002	9	11	04:50:32.860	5.2	5.2	
66.9380	-18.4560	10.0	2002	9	16	18:48:26.720	5.5	5.7	
58.3110	-31.9460	10.0	2002	10	7	20:03:54.580	4.9	5.5	
57.4490	-33.3440	10.0	2003	2	1	18:47:52.150	5.3	5.5	
71.1220	-7.5770	10.0	2003	6	19	12:59:24.410	5.6	5.0	
76.3720	23.2820	10.0	2003	7	4	07:16:44.720	5.7	5.1	
73.2730	6.4210	10.0	2003	8	30	01:04:42.340	5.0	4.8	
56.0620	111.2550	10.0	2003	9	16	11:24:52.220	5.2	5.7	
80.3140	-1.8280	10.0	2003	9	22	20:45:16.910	5.2	4.7	
50.0380	87.8130	16.0	2003	9	27	11:33:25.080	6.5	7.5	
50.0910	87.7650	10.0	2003	9	27	18:52:46.980	6.1	6.6	
50.2110	87.7210	10.0	2003	10	1	01:03:25.240	6.3	7.1	
79.1410	2.3290	10.0	2003	10	7	02:36:54.440	5.1	4.6	
74.0800	134.8210	10.0	2003	12	7	09:16:12.640	5.0	4.4	
84.4750	105.2150	10.0	2004	1	19	07:22:52.910	5.5	5.2	
71.0670	-7.7470	12.2	2004	4	14	23:07:39.940	5.8	5.6	
81.7290	119.2920	10.0	2004	6	24	22:12:37.160	4.7	-	
54.1310	-35.2590	10.0	2004	7	1	09:20:44.140	5.4	5.5	
73.8300	114.4820	10.0	2004	10	2	11:06:01.470	4.5	-	
83.2630	115.9400	10.0	2004	11	13	21:28:01.450	4.6	-	
76.1690	7.5280	10.0	2004	11	27	06:38:29.290	5.0	4.5	
84.9480	99.3100	10.0	2005	3	6	05:21:43.430	6.1	6.2	

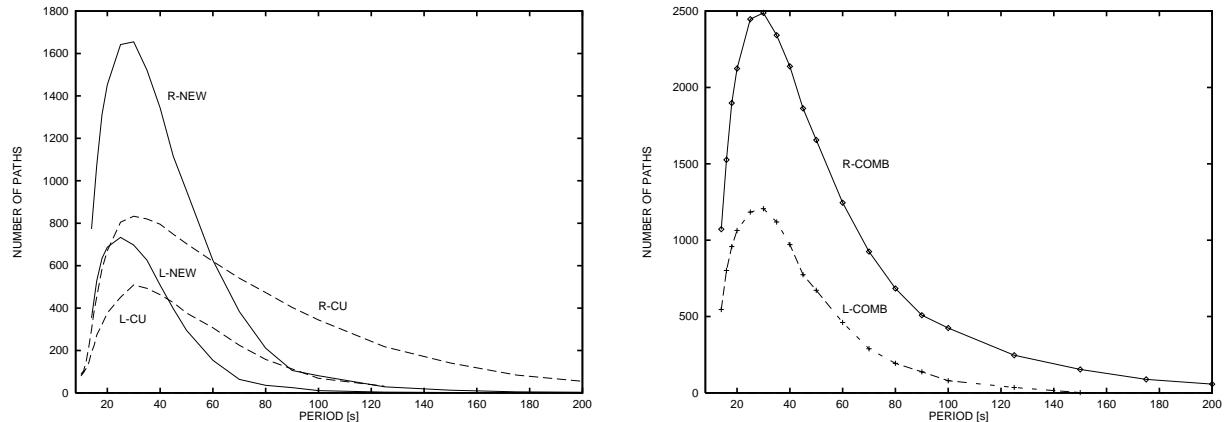


Fig. 6.3.2. The figure on the left shows the number of observed ray paths, on which Love- (L) and Rayleigh-wave (R) group velocities were measured for the preselected Colorado University data set (dashed lines, R-CU and L-CU) and during this new study (R-NEW and L-NEW). The figure on the right shows the number of ray paths in the joint data set used for the inversions in this study.

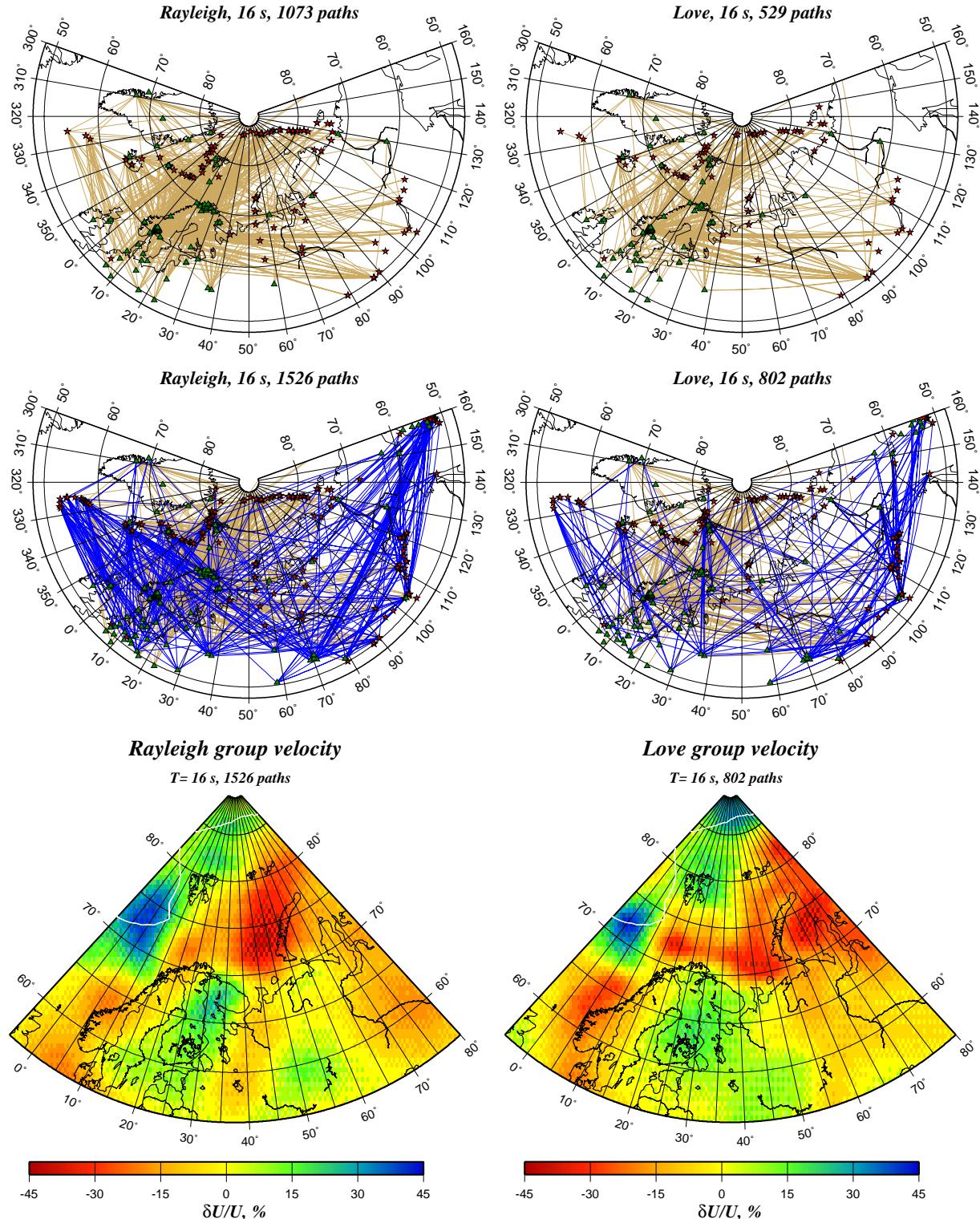


Fig. 6.3.3. Results of the 2D inversion for the group velocities of Love (on the right) and Rayleigh waves (on the left) with a period of 16 s. The two maps at the top show the ray coverage of the newly analyzed data. The two maps in the middle show the same rays plus the rays for the data from the preselected CU data set (blue). The two maps at the bottom present the 2D distribution of the inverted group velocities as deviations from the average velocity (in percent).

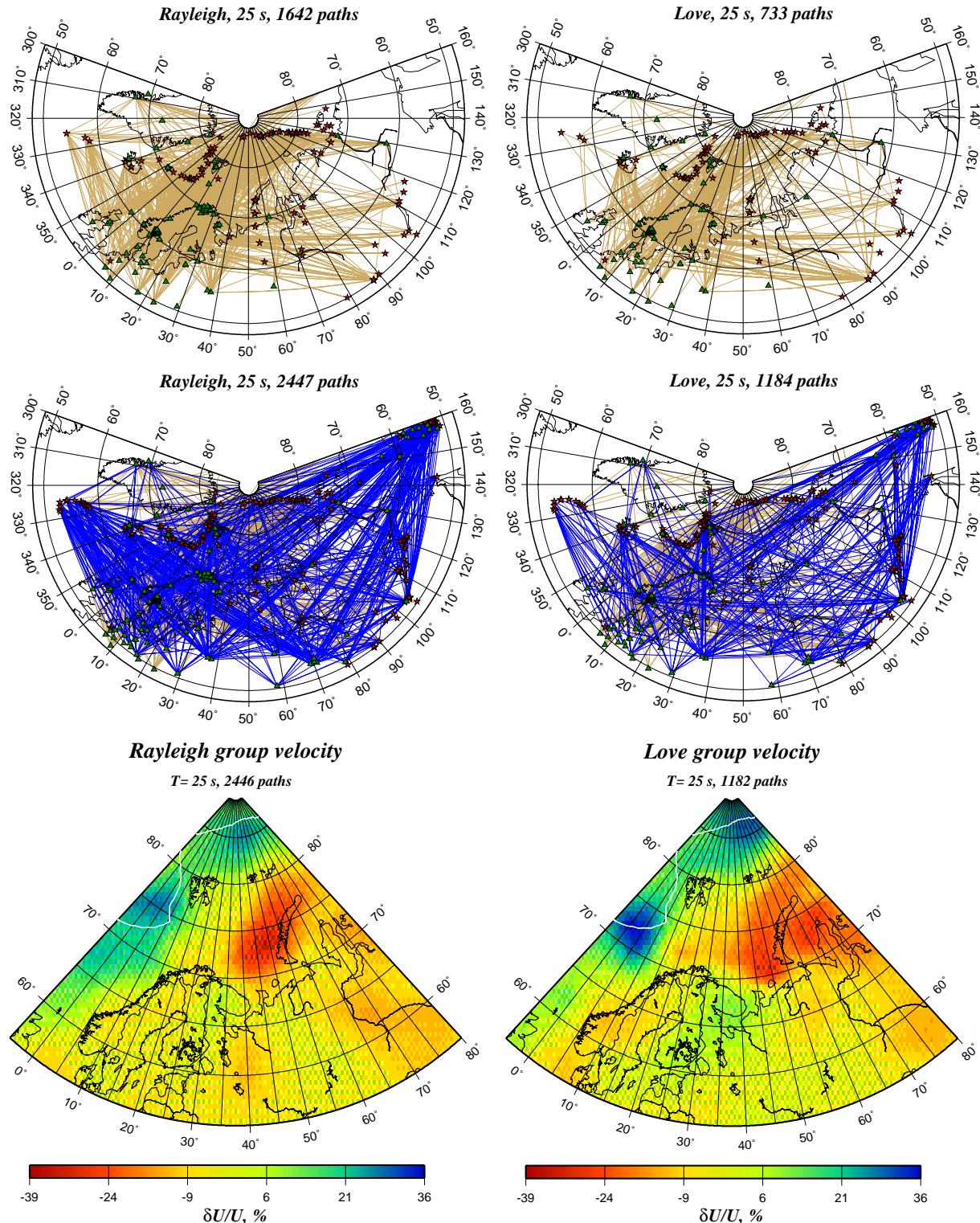


Fig. 6.3.4. Results of the 2D inversion for the group velocities of Love (on the right) and Rayleigh waves (on the left) with a period of 25 s. The two maps at the top show the ray coverage of the newly analyzed data. The two maps in the middle show the same rays plus the rays for the data from the preselected CU data set (blue). The two maps at the bottom present the 2D distribution of the inverted group velocities as deviations from the average velocity (in percent).

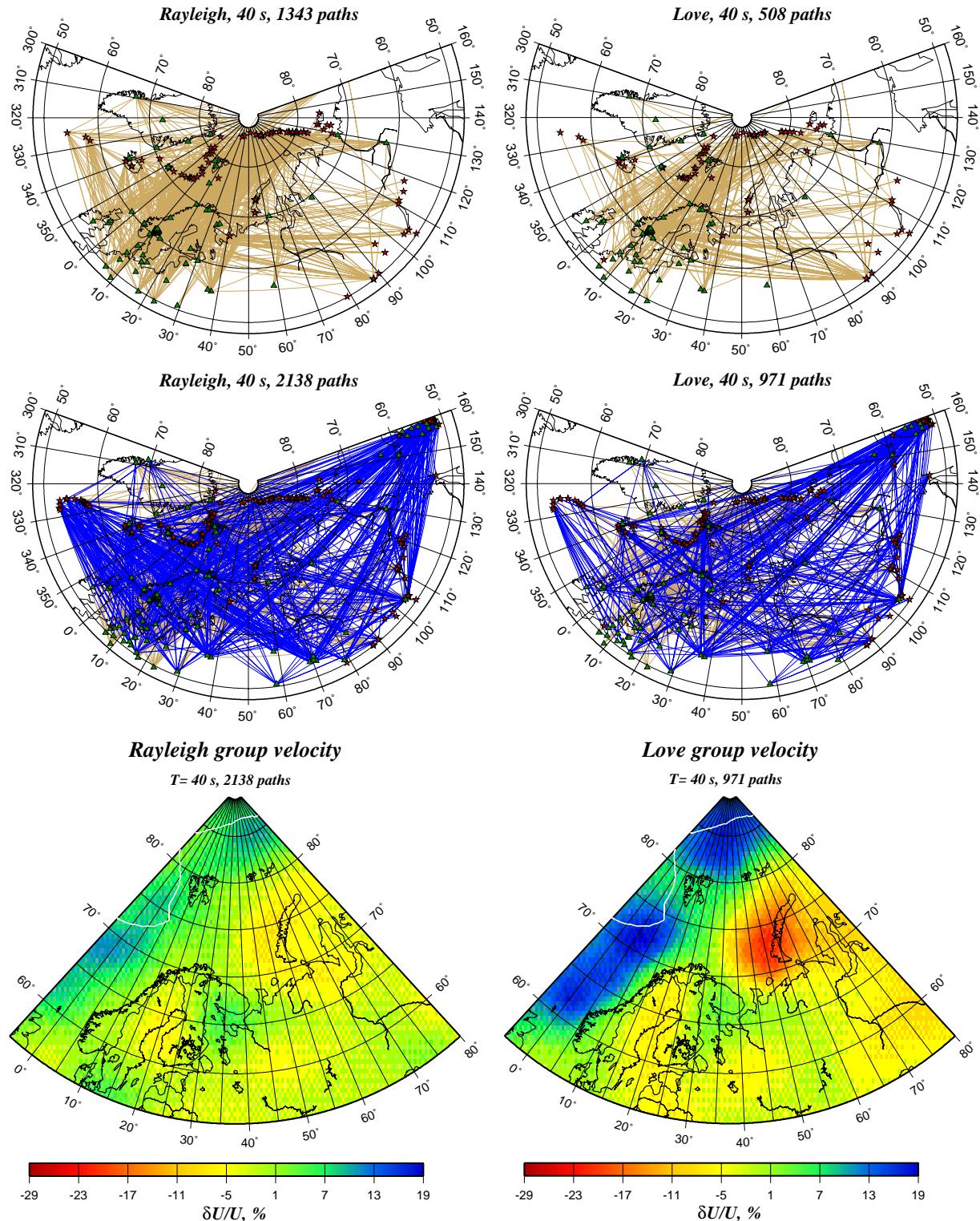


Fig. 6.3.5. Results of the 2D inversion for the group velocities of Love (on the right) and Rayleigh waves (on the left) with a period of 40 s. The two maps at the top show the ray coverage of the newly analyzed data. The two maps in the middle show the same rays plus the rays for the data from the preselected CU data set (blue). The two maps at the bottom present the 2D distribution of the inverted group velocities as deviations from the average velocity (in percent).